# Risk modeling using multiple probability distributions for the climate sensitivity

## Paul Baer, 1,2 Michael D. Mastrandrea,2 and Malte Meinshausen

<sup>1</sup> Corresponding Author: baer@stanford.edu <sup>2</sup> Center for Environmental Science and Policy, Stanford University <sup>3</sup> Swiss Federal Inst of Technology

## **Abstract**

A major challenge for climate policy is the uncertainty in the climate ensitivity, defined as the equilibrium increase of global mean surface sensitivity, defined as the equilibrium increase of global mean surface temperature in response to an equivalent doubling of CO<sub>2</sub>. The IPCC has reiterated in all its assessments the judgment that the climate sensitivity is probably between 1.5 and 4.5°C, without ever quantifying the probability that it is outside that range. The vagueness of the range has been a long-standing problem for risk modelers, who need a usable probability density function (PDF) for quantitative risk analysis.

Now, however, the problem is no longer that there is no such PDF, but Now, however, the problem is no longer that there is no such PDF, but that there are too many. At least six climate sensitivity PDFs have been published recently, which place from 3% to 30% or more of the distribution over 4.5°C. Honest appraisal makes it clear that we have today little grounds for aggregating these diverse PDFs or choosing among them; thus policy modelers and their constituencies must begin to work with multiple PDFs, and to grapple with the consequences of such multi-dimensional uncertainty.

nis poster we present a practical way of viewing multiple PDFs a their numerical characteristics, using a simple, spreadsheet-based to with a database of published climate sensitivity PDFs. Then, using a selection from the database, we show how multiple PDFs can be used in probabilistic risk models to address three different policy-relevant guestions: (1) the implied equilibrium temperature of a given level of radiative forcing (in Wm<sup>2</sup> or CO<sub>2</sub>-equivalent); (2) the implied equilibrium temperature of a specific CO<sub>2</sub> concentration given uncertainty in non-CO<sub>2</sub> forcing; and (3) the implied equilibrium temperature of a specific increase in radiative forcing, given uncertainty in current net radiative

Although this presentation focuses on examples based on global policy questions, the methods presented are of practical use for a much wider range of climate risk models, as the uncertainty in the climate sensitivity is a major component of uncertainty in all models linking emissions scenarios to impacts. We conclude with an example of how, combined with a simple impact model for potential species extinctions, multiple climate sensitivity PDFs can be used in models at a variety of scales to evaluate the extinction risks of various policy scenarios.

## Introduction and Background

Climate change as a risk management problem

- The wide range of uncertainties in the causal mechanisms of climate change (e.g., emissions, concentrations, global temperature change. change (e.g., emissions, concernations, grown emperature viewers regional change, local impacts) make it necessary to address it as a problem of risk management. Already numerous analyses have been published which project future temperature change or other impacts using probabilistic (Monte Carlo) models (e.g., [11,13,17,19,20,21])
- However, the risk management paradigm generally assumes that the probabilities of future outcomes can be known with reasonable
- For climate change, this is simply not the case; important parameters are uncertain in multiple dimensions, and can be described as "ambiguous probabilities" [5] Some previous work has addressed the issue of utilizing subjective probability from multiple experts in risk analyses (e.g., [4,10,15])
- The climate sensitivity, defined as the equilibrium response of global mean temperature to a doubling of CO<sub>2</sub>, provides a crucial case for examining these issues.

Climate sensitivity as a key uncertainty in climate risk

- It is widely recognized that the climate sensitivity is a crucial link in ion of the risk associated with future anthropogenic
- The IPCC has historically placed climate sensitivity between 1.5° and  $4.5^{\circ}$ C, with no quantification of the likelihood that it is outside that
- Consequences of the same concentrations vary enormously with climate sensitivity at the two ends of the range

probably "manageable"

- At 1.5°C, a doubling of CO<sub>2</sub> concentrations puts us slightly above today's global mean temperature not without likely harms but
- At 4.5°C, a doubling of CO<sub>2</sub> concentrations approaches the temperature increase from the last ice age to the present, with a high likelihood of catastrophic impacts.

Treating the climate sensitivity as a probability distribution (PDF)

- There has long been interest in representing climate sensitivity as a probability distribution, or PDF.
- In Morgan and Keith's 1995 expert survey, 12 of 16 experts estimated the likelihood that the climate sensitivity exceeded 4.5°C to be 5% or more, including many who estimated a 10-25% risk that it exceeded that level. This survey hasn't been repeated, but there is little reason to believe that it would show significantly smaller ranges.
- Computational methods (Bayesian and Monte Carlo) [1,6,8,12] have put from 3% to 30% or more of the distribution over 4.5°. (See Table 1)
- ', There is no "objective" way to aggregate these PDFs, nor strong justification for choosing one over all the others on the basis of scientific argument.
- Scientific argument.
  This is not to say that reasonable scientific arguments cannot be made why (for example) climate sensitivities over 8° or 10°C are much more unlikely than estimated by certain methods (e.g., [8,12]).
- Nonetheless, the complexity of the climate system and the possibility for true "surprise" makes clear the difficulty of arguing for strong constraints on the response to a large increase in radiative forcing; that is, even very high climate sensitivities can be reasonably assigned non-zero probabilities.

## Interpreting Multiple PDFs

- A PDF is a statement of belief, not a statement about the world. There is no "true" or "correct" PDF for the climate sensitivity [2].
- "Objective" computational PDFs still have irreducibly subjective
- The statement "The likelihood that the climate sensitivity is over 4.5°C is between 3% and 30%" requires careful interpretation [2].

## Introducing CSUW 1.0: The Climate Sensitivity Uncertainty Workbook

A tool for comparing climate sensitivity PDFs and facilitating their use in risk models

- Spreadsheet based tool to present multiple climate sensitivity PDFs for comparison and use in risk modeling Displays PDFs and CDFs (cumulative density functions) in graphic and tabular form
- Includes published (this version) and user-defined (next version) PDFs and CDFs as columns which can be cut-and-pasted into other risk
- Available from Paul Baer (baer@stanford.edu)

Current version includes ten climate sensitivity PDFs, as shown in Figure 1 and Table 1.

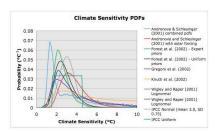


Figure 1. Graphical view of 10 climate sensitivity PDFs included in

| PDF Description  | Pct > 2    | Pct >2.5 | Pct >3 | Pct > 4 | Pct >4.5   | pct >5     |
|--|------------|----------|--------|---------|------------|------------|
| Andronova & Schlesinger (2001)     combined pdfs                           | 49%        | 41%      | 35%    | 22%     | 18%        | 14%        |
| 2 Andronove and Schlesinger (2001) with<br>solar forcing.                  | 95%        | 84%      | 67%    | 34%     | 22%        | 14%        |
| 3 Forest et al. (2002) - Expert priors                                     | 70%        | 42%      | 23%    | 68%     | 3%         | 2%         |
| 4 Forest et al. (2002) - Uniform priors<br>5 Gregory et al. (2002)         | 80%<br>79% | 65%      | 53%    | 35%     | 24%<br>29% | 19%<br>24% |
| 6 Knutti et al. (2002)   | 88%        | 79%      | 70%    | 55%     | 48%        | 42%        |
| 7 Murphy et al. (2004) weighted PDF<br>8 Wigley and Raper (2001) lognormal | 99%        | 93%      | 73%    | 27%     | 15%        | 8%         |
| based on IPCC<br>9 Normal based on IPCC, Mean 3.0°C,                       | 79%        | 55%      | 34%    | 10%     | 5%         | 3%         |
| SD 0.75°C  | 91%        | 75%      | 51%    | 10%     | 3%         | 196        |
| 10 Uniform based on IPCC   | 84%        | 67%      | 51%    | 17%     | 196        | 096        |
| Min  | 49%        | 41%      | 23%    | 68%     | 196        | 096        |
| Median   | 82%        | 66%      | 51%    | 25%     | 16%        | 1196       |
| Mean   | 81%        | 66%      | 50%    | 25%     | 17%        | 13%        |
| Max  | 99%        | 93%      | 73%    | 55%     | 48%        | 42%        |

Table 1: Summary descriptions of 10 climate sensitivity PDFs, showing the fraction of the distributions that are below or above particular threshold values of interest. PDFs highlighted in yellow are used in

## Simple risk models included in CSUW 1.0:

Likelihood for different PDFs of exceeding given temperature thresholds at equilibrium, for arbitrary stabilization levels of net forcing (in CO2-equivalent)

- Simple formula linking net forcing to equilibrium temperature does not require Monte Carlo modeling.
- ble 2 below is summary created (by hand) from 7 individual tables r different stabilization levels, using six representative PDFs.

| ppm CO2-e | 2°     | 2.5°   | 3°     | 4°     | 5°    |
|-----------|--------|--------|--------|--------|-------|
| 350       | 1-9%   | 1-4%   | 0-1%   | 0%     | 0%    |
| 400       | 8-33%  | 1-16%  | 0-10%  | 0-4%   | 0-1%  |
| 450       | 26-78% | 10-42% | 4-19%  | 0-10%  | 0-5%  |
| 500       | 43-96% | 24-74% | 11-44% | 0-16%  | 0-9%  |
| 550       | 48-99% | 40-91% | 21-69% | 6-25%  | 0-14% |
| 600       | 53-99% | 44-98% | 33-86% | 10-41% | 0-18% |
| 650       | 59-99% | 48-99% | 42-94% | 16-58% | 6-24% |

Table 2: Likelihood of exceeding various temperature thresholds at equilibrium for different levels of stabilization of radiative forcing (in ppm CO<sub>2</sub>-equivalent)

### Inverse calculation showing maximum net radiative forcing (in ppm CO2-equivalent) consistent with different probabilities of staying below various temperature thresholds.

- · Again, does not require Monte Carlo modeling.
- Table 3 below is summary for using six representative PDFs.

|      | 50%      | 67%      | 75%     | 80%     | 90%     | 95%     |
|------|----------|----------|---------|---------|---------|---------|
| Z°   | 418-574  | 401-467  | 393-449 | 387-437 | 355-410 | 339-389 |
| 2.5° | 463-689  | 440-531  | 428-506 | 420-490 | 377-452 | 356-423 |
| 3°   | 513-826  | 482-605  | 467-570 | 456-548 | 401-498 | 374-460 |
| 4º   | 629-1187 | 579-783  | 554-724 | 538-687 | 453-605 | 412-544 |
| 5°   | 771-1706 | 695-1015 | 705-920 | 680-862 | 638-735 | 455-643 |

Table 3: CO2-equivalent forcing stabilization levels associated desired probability of compliance. (Six representative PDFs; see Table 1)

#### Robust Conclusions

- A high likelihood (80-90%) of staying below 2°C increase requires stabilization on the order of 400 ppm CQ.-equivalent.
- Radiative forcing equal to a doubling of CO, has a significant likelihood (6-25%) of exceeding a 4°C temperature increase.
- These results are very similar to those reported in other literature that is either in publication [11] or in press [14].

## Using CDFs/PDFs from CSUW in external risk models

## I. Projecting equilibrium temperature, accounting for uncertainty in non-CO, radiative forcing

- Both current value and projected values of net non-CO, forcings are highly uncertain
- are highly uncertain. Well-mixed GHGs other than CO, add about 1.2 Wm² to current forcing (roughly 100 ppm CO,-equivalent). In addition to anthropogenic aerosols (estimated to equal 0 to -2 Wm² radiative forcing), other forcing agents include black carbon, tropospheric ozone, volcanic gases, changes in solar flux, and changes in land use.

A. Modeling uncertain future net non- $CO_2$  forcing

- SRES projections (marker scenarios) of net non-Co, radiative forcing in 2100 range from 0.55 Wm<sup>2</sup> to 2.45 Wm<sup>2</sup>. While the SRES scenarios do not represent a statistical sample of any process, as an indicative range, they can be represented subjectively by their mean (1.4 Wm<sup>2</sup>) and standard deviation (0.7
- A more conservative interpretation, following Wigley [20] on the assumption that there will be climate policy, is a mean of 0.75  $Wm^2$  and an SD of 0.3  $Wm^2$ .
- Using this distribution, with a the same representative selection of 6 PDFs used previously, a simple Monte Carlo model produces the results shown in Table 4 for the probability of staying below a temperature threshold of 2°C.

|               | 350 | 400 | 450 |
|---------------|-----|-----|-----|
| A&S           | 76% | 65% | 53% |
| Forest E      | 87% | 64% | 46% |
| Murphy        | 63% | 24% | 796 |
| IPCC uniform  | 75% | 44% | 24% |
| IPCC normal   | 78% | 45% | 20% |
| W&R lognormal | 83% | 58% | 35% |
| MIN           | 63% | 24% | 796 |
| MED           | 77% | 51% | 30% |
| MEAN          | 77% | 50% | 31% |
| MAX           | 87% | 65% | 53% |

Table 4. Probability of staying below 2°C above preindustrial temperature for three CO, stabilization levels and six climate sensitivity PDFs, non-CO, forcings distributed normally with mean 0.75 Wm², SD 0.3 Wm².

- B. Modeling uncertain current net non- $\mathbf{CO}_2$  forcing
- modeling uncertain current net non-CO<sub>2</sub> forcing
   Estimates of net current forcing are very uncertain. The IPCC's TAR shows a table with the uncertainty in each component, but advises against simply aggregating the uncertainties.
   In spite of this, the obvious relevance of estimating current net forcing has led several research groups to do exactly that:

   Hansen and Sato [9] add them up, treating IPCC uncertainty ranges as one-sigma normal error distributions;
   Boucher and Haywood [3] use a more sophisticated approach, performing a Monte Carlo analysis using different shape PDFs for different forcing agents.
- Knutti et al. [12] use a Bayesian computational method to estimate a PDF for net current forcing.
- These three PDFs are shown in Figure 2 below. Note that both the Boucher and Haywood and Knutti et al. PDFs have been subjectively rendered as normal distributions.
- Using these three PDFs and the same six representative climate sensitivity PDFs in a simple Monte Carlo model, estimating the risk of exceeding the 2.0°C threshold with 1 Wm² additional forcing produces the calculations shown in Table 5.

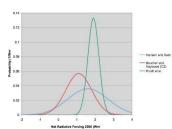


Figure 2: Three PDFs for current net radiative forcing

| Pct<2°C      | Boucher and<br>Haywood | Hansen and<br>Sato | Knutti et al. |
|--------------|------------------------|--------------------|---------------|
| A&S          | 76%                    | 67%                | 57%           |
| Forest E     | 81%                    | 67%                | 59%           |
| Murphy       | 52%                    | 35%                | 11%           |
| IPCC Uniform | 67%                    | 51%                | 33%           |
| IPCCC normal | 77%                    | 60%                | 51%           |
| Wigley       | 71%                    | 49%                | 28%           |
| MIN          | 52%                    | 35%                | 11%           |
| MED          | 73%                    | 55%                | 42%           |
| MEAN         | 71%                    | 55%                | 40%           |
| MAX          | 81%                    | 67%                | 59%           |

Table 5: Probability of exceeding 2°C increase above preindustrial temperature for 1 Wm<sup>-2</sup> additional forcing, three PDFs for current net forcing.

#### II. Projecting impacts at equilibrium

- The probabilistic output of an equilibrium temperature model using one or more of the climate sensitivity PDFs in CSUW can straightforwardly be used to project risks of impacts of various brose
- In the placeholder example below, equilibrium global temperature increase drives a probabilistic species loss function.
- The curves shown in Figure 3 indicate minimum, median and maximum estimates of the fraction of species lost (shown on the Y axis) for a given equilibrium temperature increase.
- The minimum, median and maximum loss functions are arbitrarily defined as quadratic functions of equilibrium temperature increase. Note that the figures shown below are much lower than the estimates given in [18].
- esuminates given in [15]. The functions below were entered into a simple Monte Carlo model in which, in each "rum" of the model, a value of equilibrium temperature is estimated first, and the nate in impact is calculated probabilistically using a triangular distribution using the minimum median and maximum impact values for that temperature. Table 6

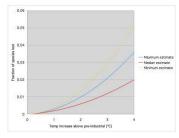


Figure 3. Placeholder example of probabilistic impacts of global mean temperature increase on species loss: min, median and mas fraction of species lost at a given temperature increase.

|     | Median     | 90th pctile | 95th pctile | Maximum    |
|-----|------------|-------------|-------------|------------|
| 400 | 0.62-0.63% | 0.97-1.00%  | 1.07-1.11%  | 1.53-1.83% |
| 450 | 1.02-1.06% | 1.63-1.71%  | 1.80-1.93%  | 2.60-3.03% |
| 500 | 1.46-1.54% | 2.34-2.46%  | 2.63-2.75%  | 3.93-4.64% |
| 550 | 1.92-2.01% | 3.16-3.29%  | 3.55-3.73%  | 5.10-6.32% |
| 600 | 2 42-2 48% | 3 96-4 16%  | 4 47-4 59%  | 6 40.8 24% |

Table 6: Range of results for placeholder species loss model, showing range of projected percent species loss (median, 90th pctile, 95th pctile, max) for six selected PDFs and five stabilization levels (in ppm CO,-equivalent)

## Interpreting the results

- In this particular example, the fairly compressed range for each "entry" (typically less than 5% for all except the maximum value) shows that the estimated impacts are not very sensitive to the climate sensitivity PDF or PDFs that are used.
- Using this table, one can give policy recommendations of the following
  - if, using these assumptions, one wishes to have a 90% chance of keeping species loss below 2.5%, atmospheric stabilization has to be no higher than about 500 ppm CO,-equivalent.
- If atmospheric CO, concentrations stabilize at 600 ppm CO2-equivalent or higher, there is at least a 10% chance that there will be a 4% loss of species.
- ue a 4% loss of species.

  The model above is highly generalizable; as long as the impacts can be specified probabilistically as a function of equilibrium temperature increase, the same basic model structure can be used. It is fairly straightforward to modify such a model to use transient temperature increase (e.g., temperature increase at 2100). So far, however, there are no simple but robust probabilistic models of transient temperature increase in the literature.

- There is not, and will not soon be, a single consensual PDF for the climate sensitivity.
- climate sensitivity.

  This is a general property of subjective probability estimates of uncertain parameters, including (for example) the rate of ocean heat uptake, the future strength of the carbon sink, etc.

  Accordingly, it will never be possible to say something like "the probability of temperature exceeding 2"c in 2100 is 3%, if CO, emissions are held to 400 GiC over the 21st century" the result of proper modelling, using multiple PDFs will always be a range, even if (as shown in the species loss model above) it is a small range.
- The design of policies which set risk targets must therefore not only specify an acceptable risk of exceeding some threshold, they must address how such a target will be set given the range of relevant input PDPs.

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